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Multiple solutions for the Riemann problem in the Porous Shallow Water Equations

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Abstract

The Porous Shallow water Equations are widely used in the context of urban flooding simulation. In these equations, the solid obstacles are implicitly taken into account by averaging the classic Shallow water Equations on a control volume containing the fluid phase and the obstacles. Numerical models for the approximate solution of these equations are usually based on the approximate calculation of the Riemann fluxes at the interface between cells. In the present paper, it is presented the exact solution of the one-dimensional Riemann problem over the dry bed, and it is shown that the solution always exists, but there are initial conditions for which it is not unique. The non-uniqueness of the Riemann problem solution opens interesting questions about which is the physically congruent wave configuration in the case of solution multiplicity.

1 Introduction

The Shallow water Equations are widely used to study urban flooding events. These hyperbolic partial differential equations are non-linear, and admit the existence of discontinuities (moving bores, standing hydraulic jumps, wetting-drying boundaries) that develop in finite time, also if the initial conditions are smooth, i.e. continuous with higher-order derivatives [1, 2]. The Riemann problem is a special initial value problem where the initial conditions are discontinuous, and it is used to model the cited flow field discontinuities [1, 2].

Modern numerical models for the simulation of flooding phenomena are often based on the approximate or exact solution of the classic Riemann problem for the Shallow water Equations, which always exists and it is unique [1]. The solution of the Riemann problem on dry bed [3] has started the modern approaches in the evaluation of flooded areas extent. This solution, which inherits the uniqueness property from the general Riemann problem, exhibits a very simple structure, because it consists of a single rarefaction wave that connects the left state with the right dry bed [1, 3].

In the urban environment, the numerical calculations are very time consuming, because the modelling of building geometric details requires fine computational grids [4]. In order to reduce the computational burden, an alternative approach based on the Porous Shallow water Equations [5, 6, 7] has been introduced. In this approach, which corresponds to the averaging methods used for multiphase flows [8], the solid obstacles are implicitly taken into account by averaging the classic Shallow water Equations on a Representative Elementary Volume containing the fluid phase and the obstacles [5, 7]. In the single-porosity models, which are the object of the present paper, the fraction of the unit-surface area that is not occupied by obstacles is called porosity [5, 6, 7]. It is possible to demonstrate that the Porous Shallow water Equations are non-linear and hyperbolic [6, 9], and then may exhibit discontinuities. Moreover, a rotational invariance property is satisfied [6].

In realistic applications, the porosity can be discontinuous, because the fraction of the urban areas occupied by buildings can be strongly variable in space, and this leads to the open question about the influence of geometric discontinuities on the solution of the Riemann problem for the Porous Shallow water Equations. Despite the existence of some example solution (see, for example, [6, 9, 10]), there is not a systematic study of the general Riemann problem for the Porous Shallow water Equations.

It is possible to demonstrate [11] that the solution of the dam-break problem, i.e. the particular case of Riemann problem with initial zero velocity, depends on the definition of the geometric-discontinuity inner-structure. In particular, most of the existing geometric discontinuity definitions, such as those contained in [6] and [10], may lead to unphysical solutions (increase of the flow energy through the discontinuity). Other definitions, such as those by [12] and [13], are either internally inconsistent or imply the incompleteness of the problem solution. Interestingly, the solution of the dam-break problem always exists and it is unique if an appropriate parameterization of the stationary weak solutions of the Porous Shallow water Equations is used to supply the inner structure of the geometric discontinuity, as made in [11].

In the present paper, the exact solution of the one-dimensional Riemann problem for the Porous Shallow water Equations is presented, assuming that the right state is initially dry and that the jump definition is the same of [11]. This special Riemann problem on dry bed seems a good candidate in order to explore interesting problems encountered when the flooding of urban areas with strong porosity gradients and initial dry bed is modelled.

2 The position of the Riemann problem for the Porous Shallow water Equations

The rotational invariance of the two-dimensional Porous Shallow water Equations allows considering the one-dimensional framework

$$\frac{\partial \phi \mathbf{u}}{\partial t} + \frac{\partial \phi \mathbf{f}(\mathbf{u})}{\partial x} + \mathbf{h}(\mathbf{u}) \frac{\partial \phi}{\partial x} = 0, \quad (1)$$

where the following discontinuous initial conditions are considered:

$$\mathbf{u}(x, 0) = \begin{cases} \mathbf{u}_L, & x < 0 \\ \mathbf{u}_R, & x > 0 \end{cases}, \quad \phi(x) = \begin{cases} \phi_L, & x < 0 \\ \phi_R, & x > 0 \end{cases}. \quad (2)$$

In Eqs. (1)-(2), the symbols are defined as follows: x is the space variable, t is the time variable, $h(x, t)$ is the water depth, $u(x, t)$ is the depth-averaged velocity, $\mathbf{u} = (h \quad hu)^T$ is the vector of the flow conserved variables, $\mathbf{f}(\mathbf{u}) = (hu \quad hu^2 + gh^2/2)^T$ is the vector of the fluxes, $\mathbf{h}(\mathbf{u}) = (0 \quad -gh^2/2)^T$ is a vector that takes into account the porosity gradients effect, $\mathbf{u}_L = (h_L \quad h_L u_L)^T$ and $\mathbf{u}_R = (h_R \quad h_R u_R)^T$ are the initial left and right flow states, while $\phi(x)$ is the porosity. The aspect ratio AR is defined as the ratio between the right and the left porosity, that is $AR = \phi_R / \phi_L$.

In order to consider the Riemann problem on dry bed, in the following it will be assumed that the state \mathbf{u}_R coincides with the dry bed state $\mathbf{0} = (0 \quad 0)^T$. This implies that the flow through the geometric discontinuity is always directed from left to right.

2.1 The porosity discontinuity

The porosity discontinuity induces a standing flow discontinuity at $x = 0$ for $t > 0$. This standing discontinuity, where the non-conservative product $h^2 \partial \phi / \partial x$ is active, is not autonomous [14], and the missing physical information must be supplemented. Congruently with [11], it is assumed here that the states $\mathbf{u}_1 = (h_1 \quad h_1 u_1)^T$ and $\mathbf{u}_2 = (h_2 \quad h_2 u_2)^T$, immediately to the left ($x = 0^-$) and to the right ($x = 0^+$) of the geometric discontinuity in $x = 0$, are connected by a stationary solution of Eq. (1). This means that the flow energy is conserved through the geometric discontinuity, unless a hydraulic jump is present. In addition, it is assumed that ϕ varies monotonically through the porosity jump. This means that critical state conditions can be attained only at the left ($AR \geq 1$) or at the right ($AR < 1$) of the geometric discontinuity. In the following $F_1 = u_1 / \sqrt{gh_1}$ and $F_2 = u_2 / \sqrt{gh_2}$ are the Froude numbers corresponding to the states \mathbf{u}_1 and \mathbf{u}_2 , respectively.

2.2 Ritter solution ($AR = 1$)

When the porosity discontinuity is absent ($AR = 1$), the solution of the Riemann problem on the dry bed consists of a unique rarefaction wave contained in the first characteristic field that connects the left state \mathbf{u}_L to the dry bed state $\mathbf{u}_R = \mathbf{0}$, while $\mathbf{u}_1 = \mathbf{u}_2$. Along the rarefaction, the velocity and the fluid depth are constrained by [1, 2]

$$u + 2\sqrt{gh} = u_L + 2\sqrt{gh_L} \quad (3)$$

and

$$x = \lambda_1(\mathbf{u})t, \quad (4)$$

where $\lambda_1(\mathbf{u}) = u - \sqrt{gh}$. In particular, the left and the right edges of the rarefaction wave have speed $\sigma_L = u_L - \sqrt{gh_L}$ and $\sigma_R = u_L + 2\sqrt{gh_L}$, respectively, with $\sigma_L < \sigma_R$. The velocity u increases (

$\partial u/\partial x > 0$) along the rarefaction from the left edge ($u = u_L$) to the right edge ($u = u_L + 2\sqrt{gh_L}$), while the flow depth decreases ($\partial h/\partial x < 0$) from the left edge ($h = h_L$) to the right edge ($h = 0$). If $F_L = u_L/\sqrt{gh_L}$ is the Froude number corresponding to the left state \mathbf{u}_L , it is easy to see that the Froude number $F = u/\sqrt{gh}$ increases continuously from the value $F = F_L$ at the left edge, and tends to infinity at the right edge.

The Froude number F_L rules the following three cases:

Case Ra. When $F_L \leq -2$, the rarefaction wave develops entirely to the left of $x = 0$ because σ_R is negative, and $u < 0$ everywhere. In this case, $\mathbf{u}_1 = \mathbf{0}$.

Case Rb. When $-2 < F_L \leq 1$, the wave develops through $x = 0$ because $\sigma_L \leq 0$ and $\sigma_R \geq 0$. The eq. (4) implies that critical flow conditions are obtained in $x = 0$, with flow velocity $u = \sqrt{gh}$ directed from left to right. In this case $\mathbf{u}_1 = \left(h_L (F_L + 2)^2 / 9 \quad \sqrt{gh_L^3} (F_L + 2)^3 / 27 \right)^T$.

Case Rc. When $F_L > 1$, the rarefaction wave develops entirely to the right of $x = 0$ because σ_L is positive. This implies that $\mathbf{u}_1 = \mathbf{u}_L$ for $x \leq 0$. In addition, the velocity u is positive everywhere, and the flow is supercritical.

2.3 Energy limit

When $u_1 > 0$, a necessary condition for the flow to be smooth through the geometric discontinuity is that the energy of the incoming flow \mathbf{u}_1 is not less than the minimum energy (critical state) required at the outlet. This implies that the constraint

$$2 + F_1^2 \geq 3AR^{-2/3}F_1^{2/3}, \quad (5)$$

is satisfied. From the study of eq. (5) it follows that:

(a) $AR \geq 1$ is required when $F_1 = 1$, and this means that the critical flow cannot freely pass through a porosity decrement.

(b) The eq. (5) is satisfied by any value of F_1 for $AR \geq 1$. In other words, any flow is able to pass through a porosity increment.

(c) When $AR < 1$, the eq. (5) is satisfied by the values $F_1 \in]0, F_{sub}(AR)] \cup [F_{sup}(AR), \infty[$, where $F_{sub}(AR)$ and $F_{sup}(AR)$ are two limits depending on AR , with $0 < F_{sub}(AR) < 1$ and $F_{sup}(AR) > 1$. The values $F_1 = F_{sub}(AR)$ or $F_1 = F_{sup}(AR)$ at the inlet of the porosity discontinuity satisfy the eq. (5) with the equality sign, and imply $F_2 = 1$ at the outlet of the porosity decrement.

2.4 Hydraulic jump limit for supercritical flows

When $AR < 1$, it is interesting to consider the case where critical flow conditions are obtained at the exit of the discontinuity ($F_2 = 1$), while a subcritical stationary flow profile develops through the geometric discontinuity and $F_1 = F_{sub}(AR)$. The supercritical flow $F_{jump}(AR)$ connected by a standing hydraulic jump to the subcritical flow at the inlet of the geometric discontinuity satisfies the limit condition

$$F_{jump}(AR) = F_{sub}(AR) \left(\frac{2}{\sqrt{1+8F_{sub}^2(AR)}-3} \right)^{3/2}. \quad (6)$$

The supercritical flow with Froude number less than $F_{jump}(AR)$ is not able to sustain a hydraulic jump at the inlet of the sudden porosity decrement, and a shock moving upstream may be obtained.

2.5 Hydraulic hysteresis region

If the equality sign and $F_L = F_1$ are taken in eq. (5), the limit condition

$$AR = F_L \left(\frac{3}{2} \right)^{3/2} \left(1 + \frac{F_L^2}{2} \right)^{-3/2}, \quad (7)$$

is obtained [15]. The eq. (7) is represented in Figure 1 (continuous black line), and the corresponding graph consists of a subcritical reach AB and a supercritical reach BC. Based on eq. (5), the points of the plane (F_L, AR) above the energy limit curve of eq. (7) refer to flows \mathbf{u}_L that are able to pass through the geometric discontinuity remaining smooth.

If $F_L = F_{jump}(AR)$ is taken, the eq. (6) can be solved with respect to $F_{sub}(AR)$. Recalling that $F_1 = F_{sub}(AR)$ satisfies eq. (5) with the equality sign, the following limit condition [16, 17] is obtained

$$AR = \left(\frac{3}{2} \right)^{3/2} F_L \left(1 + \frac{F_L^2}{2} - \frac{\left(\sqrt{1+8F_L^2} - 3 \right)^3}{16 \left(\sqrt{1+8F_L^2} - 1 \right)} \right)^{-3/2}. \quad (8)$$

The corresponding curve, which is plotted in Figure 1 (curve BD with dashed black line), lies above the energy limit curve ABC. Based on eq. (6), the points of the plane (F_L, AR) below the standing hydraulic jump limit BD are referred to supercritical flows \mathbf{u}_L that do not pass through the geometric discontinuity, because they are turned into subcritical states \mathbf{u}_1 by a backward moving shock [16, 17].

The region of the plane between the curves BC and BD admits both an upstream supercritical flow passing through the discontinuity (above BC) and a supercritical upstream flow that is turned in subcritical (below BD). This phenomenon is called hydraulic hysteresis [16, 17]. Interestingly, in the hysteresis region, a third flow condition is possible, namely a standing hydraulic jump through the geometric discontinuity, where the upstream supercritical flow is turned in subcritical [17].

3 Solution configurations for the Riemann problem with right dry state

Depending on the left Froude number F_L and on the aspect ratio AR , eight different solution configurations for the Riemann problem on dry bed are possible. These solution configurations, whose fields of existence are plotted in Figure 1, are listed as follows.

3.1 Solution configuration not depending on AR

SC (1). In this solution configuration, the condition $F_L \leq -2$ is satisfied, and the solution configuration coincides with the *Case Ra* of subsection 2.2. A rarefaction wave develops entirely to the left of the geometric discontinuity, and the velocity u is non-positive everywhere. In other words, there is not interaction between the flow and the geometric discontinuity.

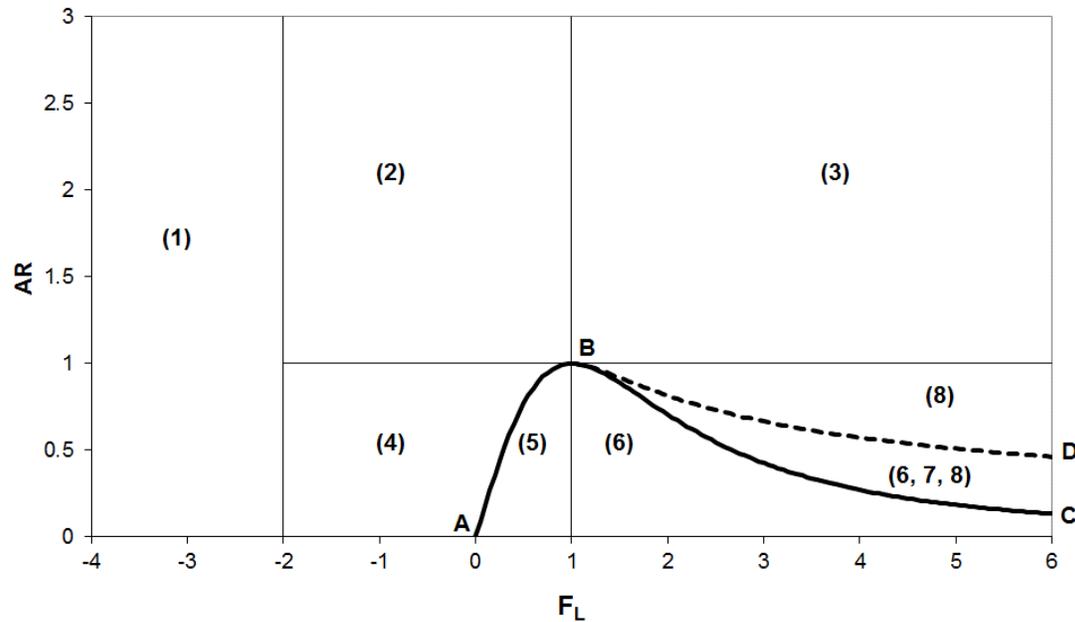


Figure 1: Field of existence of the solution configurations.

3.2 Solution configurations with $AR \geq 1$ (sudden porosity increment)

In this group of solution configurations, the flow corresponding to the Ritter solution (*Case Rb* and *Case Rc* of subsection 2.2) immediately to the left of $x = 0$ encounters a sudden porosity increment. For this reason, the geometric discontinuity is not an obstacle to the flow.

SC (2). In this solution configuration, $-2 < F_L \leq 1$ (*Case Rb*). A rarefaction connects the state \mathbf{u}_L to the critical state \mathbf{u}_1 , and the energy of the flow is sufficient to pass through the geometric discontinuity because $F_1 = 1$ satisfies the eq. (5) with $AR \geq 1$. The supercritical state \mathbf{u}_2 is connected to \mathbf{u}_1 by means of the conditions of energy and discharge conservation. Finally, the dry bed state \mathbf{u}_R is connected to \mathbf{u}_2 by means of a rarefaction.

SC (3). In this solution configuration, $F_L > 1$ (*Case Rc*). The state \mathbf{u}_1 coincides with the supercritical state \mathbf{u}_L , and the energy of the flow is sufficient to pass through the geometric discontinuity because $F_1 > 1$ satisfies the eq. (5) with $AR \geq 1$. The supercritical state \mathbf{u}_2 is connected to \mathbf{u}_1 by means of the conditions of energy and discharge conservation, while a rarefaction connects the state \mathbf{u}_2 to the dry bed.

3.3 Solution configurations with $AR < 1$ (sudden porosity decrement)

SC (4). In this solution configuration, $-2 < F_L \leq F_{sub}(AR)$ (*Case Rb*). A rarefaction connects the state \mathbf{u}_L to the subcritical state \mathbf{u}_1 , which is characterized by $F_1 = F_{sub}(AR)$. The state \mathbf{u}_2 is critical, and it is connected to \mathbf{u}_1 by means of the conditions of energy and discharge conservation. Finally, the state state \mathbf{u}_2 is connected to the dry bed by means of a rarefaction.

SC (5). In this solution configuration, $F_{sub}(AR) < F_L \leq 1$ (*Case Rb*). A shock moving upstream connects the subcritical (or critical) state \mathbf{u}_L to the subcritical state \mathbf{u}_1 , which is characterized by $F_1 = F_{sub}(AR)$. The state \mathbf{u}_2 is critical, and it is connected to \mathbf{u}_1 by means of the conditions of energy and discharge conservation. Finally, the state state \mathbf{u}_2 is connected to the dry bed by means of a rarefaction.

SC (6). In this solution configuration, $1 < F_L \leq F_{jump}(AR)$ (*Case Rc*). A shock moving upstream (if $1 < F_L < F_{jump}(AR)$) or standing in $x = 0^-$ (if $F_L = F_{jump}(AR)$) connects the supercritical state \mathbf{u}_L to the subcritical state \mathbf{u}_1 , which is characterized by $F_1 = F_{sub}(AR)$. The state \mathbf{u}_2 is critical, and it is connected to \mathbf{u}_1 by means of the conditions of energy and discharge conservation. Finally, the critical state \mathbf{u}_2 is connected to the dry bed by means of a rarefaction.

SC (7). In this solution configuration, $F_{sup}(AR) \leq F_L \leq F_{jump}(AR)$ (*Case Rc*). The state \mathbf{u}_1 coincides with the supercritical state \mathbf{u}_L , while the critical state \mathbf{u}_2 is connected to \mathbf{u}_1 by means of the conditions of discharge conservation only, and a hydraulic jump is located through the geometric discontinuity. Finally, the critical state \mathbf{u}_2 is connected to the dry bed by means of a rarefaction.

SC (8). In this solution configuration, $F_{sup}(AR) \leq F_L$ (*Case Rc*), and the flow is free to pass through the porosity decrement. The state \mathbf{u}_1 coincides with the supercritical state \mathbf{u}_L . The state \mathbf{u}_2 is connected to the state \mathbf{u}_1 by means of the conditions of energy and discharge conservation. A rarefaction wave, which connects the supercritical state \mathbf{u}_2 to the dry bed, develops entirely on the right of the geometric discontinuity.

4 Discussion

The solution of the Riemann problem on the dry bed for the classic Shallow water Equations is unique, and consists of a rarefaction wave connecting the left state to the right dry bed. Conversely, the Riemann problem on the dry bed for the Porous Shallow water Equations may exhibit numerous solution configurations, which differ for the type of waves, as well their number and position. The Table 1 resumes these solution configurations, described in the Section 3, using the following symbols: R = rarefaction wave, SW = standing wave (geometric discontinuity), S = shock.

The direct calculations show that the solution of the Riemann problem always exists for given flow initial conditions and geometric discontinuity characteristics. Nonetheless, there are cases where this solution is not unique (bifurcation phenomenon), and the inspection of Figure 1 confirms that there is a field of the plane (F_L, AR) where three distinct solution configurations (*SC* (6), *SC* (7), and *SC* (8)) are possible for given initial conditions and geometric discontinuity characteristics. From a physical point of view, this field coincides with the case where the supercritical incoming flow has energy greater than that strictly required to pass through the porosity reduction, and this energy can be dissipated through a standing hydraulic jump or a moving shock. These solutions are the unsteady flow counterparts of the multiple steady states in converging channels [16, 17].

The non-uniqueness of the Riemann problem solution is expected in the case that geometric discontinuities are present, and it is well documented in the context of the Shallow water Equations with a bed step [18, 19].

Solution configuration	Waves
SC (1)	R
SC (2)	R, SW, R
SC (3)	SW, R
SC (4)	R, SW, R
SC (5)	S, SW, R
SC (6)	S, SW, R
SC (7)	SW, R
SC (8)	SW, R

Table 1: Solution configurations

5 Conclusions

The solution of the Riemann problem over the dry bed for the Porous Shallow water Equations always exists, and the main parameters on which it depends are the initial left Froude number F_L and the aspect ratio AR , i.e. the ratio between the right and the left porosity. In addition, it is demonstrated that, contrarily to the case of the classic Shallow water Equations (Ritter solution), numerous different wave configurations are possible for $t > 0$, despite of the very simple initial conditions. These flow configurations differ for the number and the type of waves exhibited by the solution.

Interestingly, there are values of F_L and AR for which the problem solution is not unique, and multiple flow configurations are possible. This phenomenon, which appears for supercritical left states, has never been documented in the context of the Porous Shallow water Equations. The non-uniqueness of the solutions opens interesting questions about which is the physically congruent wave configuration in the case of solution multiplicity, and about the foundations of the Porous Shallow water Equations themselves.

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